

# Dynamics of Metal–Silicate Equilibration in a Terrestrial Magma Ocean

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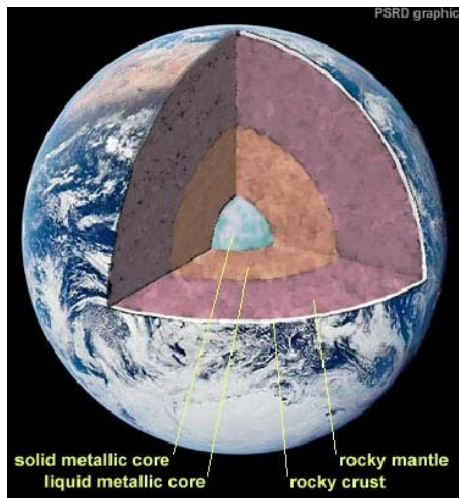
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# Outline

- Motivation
- Introduction - core formation and the early Earth
- Problem formulation
- Numerical realization
- Experiments and key parameters
- Prospects

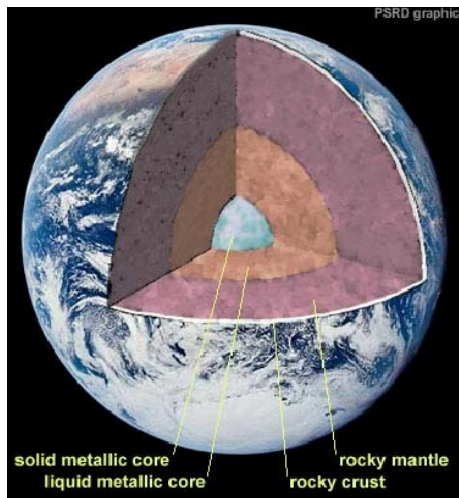
# Motivation



- Understanding of the process of core formation
- The metal droplet scenario

(Stevenson, Origin of the Earth, 1990)

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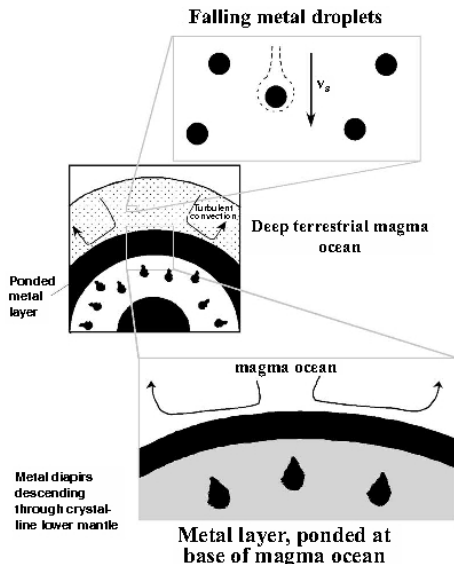
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# Primitive Earth

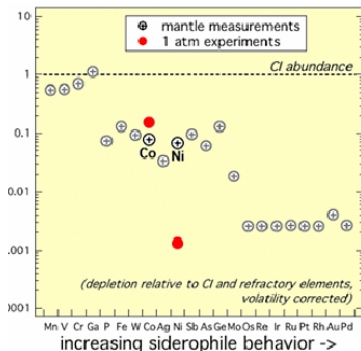
- Accreted from chondritic material
- Magma oceans were episodically formed
- Early differentiation involved the separation of silicate and Fe-rich metal
- Magma ocean condition
  - pressure 30 – 60 *GPa*
  - temperature > 2200 *K*

(Chabot, *Geochimica et Cosmoch.*, 2005)

# Core formation scenario



# Chemical constraints



**Figure:** Concentration of siderophile elements in Earth's mantle.

(Taylor, Planet. Sc. Research, 2005)

# Formulation - The model

## Basic assumptions

- Coexistence of two immiscible liquids of Fe–alloy and silicate
- Stationary situation  $\Rightarrow$  no acceleration
- Axisymmetrical problem  $\Rightarrow$  two independent variables  $(\theta, r)$



# Formulation - The model

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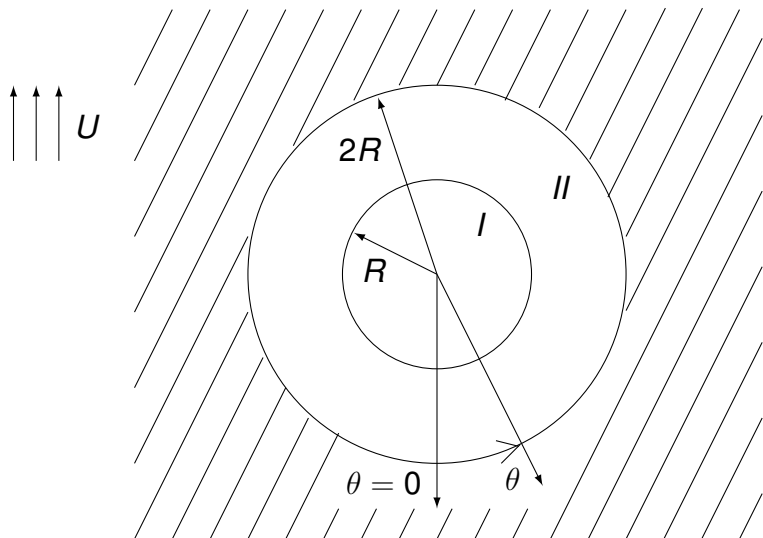
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# Geometry of a falling drop



# Basic equation

## Conservation equations for chemical exchange

- in metal

$$\frac{\partial C^I}{\partial t} = -\nabla \cdot \mathbf{F}^I = \nabla \cdot (D_{Fe} \nabla C^I - \mathbf{v}^I C^I)$$

- in silicates

$$\frac{\partial C^{II}}{\partial t} = -\nabla \cdot \mathbf{F}^{II} = \nabla \cdot (D_{Si} \nabla C^{II} - \mathbf{v}^{II} C^{II})$$

# Velocity field

## Conservation equations

- in metal

$$\nabla \cdot \mathbf{v}^I = 0$$

$$-\nabla P^I + \mu_{Fe} \nabla^2 \mathbf{v}^I = \mathbf{0}$$

- in silicates

$$\nabla \cdot \mathbf{v}^{II} = 0$$

$$-\nabla P^{II} + \mu_{Si} \nabla^2 \mathbf{v}^{II} = \mathbf{0}$$

## Boundary condition

$$\mathbf{v}_r^{II} = \mathbf{v}_r^I = 0$$

$$\mathbf{v}_\theta^{II} = \mathbf{v}_\theta^I$$

$$\sigma_{r\theta}^{II} = \sigma_{r\theta}^I$$

# Velocity field

## In the blob

$$v_r^I = \left(1 - \frac{r^2}{R^2}\right) \frac{U\mu^E \cos \theta}{2(\mu^I + \mu^E)}$$

$$v_\theta^I = \left(\frac{2r^2}{R^2} - 1\right) \frac{U\mu^E \sin \theta}{2(\mu^I + \mu^E)}$$

## In the magma

$$v_r^{II} = \left(-1 - \frac{\mu^I}{2(\mu^I + \mu^{II})} \frac{R^3}{r^3} + \frac{2\mu^I + 3\mu^{II}}{2(\mu^{II} + \mu^I)} \frac{R}{r}\right) U \cos \theta$$

$$v_\theta^{II} = \left(1 - \frac{\mu^I}{4(\mu^I + \mu^{II})} \frac{R^3}{r^3} - \frac{2\mu^I + 3\mu^{II}}{4(\mu^{II} + \mu^I)} \frac{R}{r}\right) U \sin \theta$$

## Falling velocity

$$U = \frac{2(\rho^E - \rho^I)gR^2}{3\mu^E} \frac{\mu^E + \mu^I}{2\mu^E + 3\mu^I}$$

# Streamlines

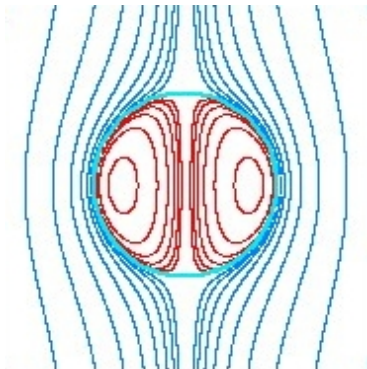


Figure: Streamlines for steady state.

# Boundary Condition

- Interface

$$[\mathbf{F} \cdot \mathbf{n}]_{-}^{+} = 0 \rightarrow \left[ D \frac{\partial C}{\partial r} \right]_{-}^{+} = 0$$

$$K = \frac{C'}{C''}$$

- External boundary

$$\theta \in (0, \frac{\pi}{2}) \rightarrow C(t) = \text{const}$$

$$\theta \in (\frac{\pi}{2}, \pi) \rightarrow (\mathbf{v} \cdot \nabla C) = 0$$



# Time approach - ADI

- Based on Crank–Nicolson scheme with RHS

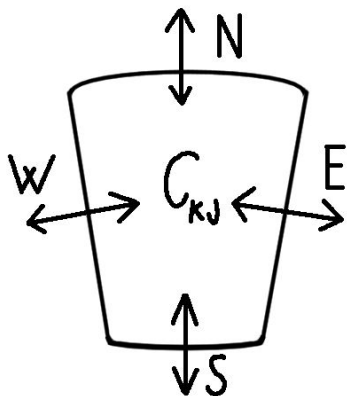
$$\partial_t \mathbf{C} = \Lambda \mathbf{C} = (\Lambda_r + \Lambda_\theta) \mathbf{C}$$

- Solved by Peaceman–Rachford scheme

$$\left(1 - \frac{\Delta t}{2} \Lambda_r\right) \mathbf{C}^{n+\frac{1}{2}} = \left(1 + \frac{\Delta t}{2} \Lambda_\theta\right) \mathbf{C}^n$$

$$\left(1 - \frac{\Delta t}{2} \Lambda_\theta\right) \mathbf{C}^{n+1} = \left(1 + \frac{\Delta t}{2} \Lambda_r\right) \mathbf{C}^{n+\frac{1}{2}}$$

# Spatial approach – finite volume formulation



## Time shots

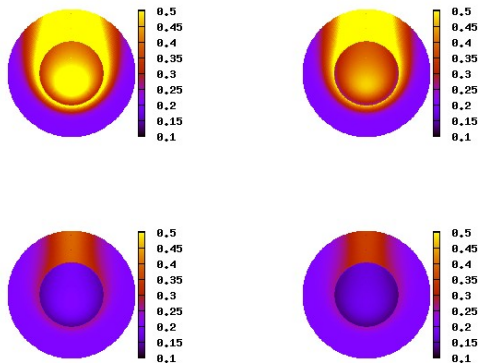


Figure:  $u_{int} = 1.$ ,  $u_{ext} = 0.4$ ,  $u_{out} = 0.2$ ,  $K = 0.5$ ,  $H_C = 0$ .

## Experiments

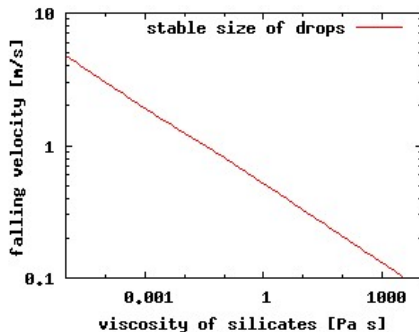
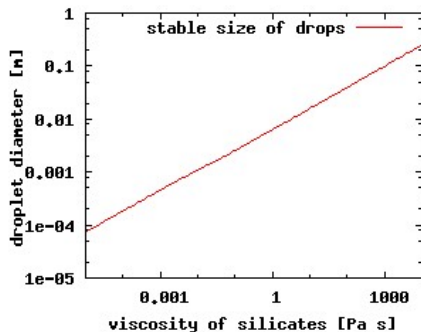
- 1  $K = \text{const}$   
(in the beginning the drop is not in equilibrium)
- 2  $K = K(\text{depth})$

## Experiments

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# Stable size of droplets

- $We = \frac{\rho^E R U^2}{2\sigma}$
- $We_c = 4\pi(1 + \frac{1}{S})$  ,  $S = \frac{\rho^l}{\rho^E}$



# Experiment nb<sup>o</sup> 1 – Dimensionless form

- $r = Rr^*$
- $\mathbf{v} = U\mathbf{v}^*$
- $C = C_0 C^*$
- $t = t_0 t^* = \frac{R}{U} t^*$

$$\frac{\partial C^*(r^*, \theta^*, t^*)}{\partial t^*} = \frac{1}{Pe} \nabla^2 C^*(r^*, \theta^*, t^*) - \mathbf{v}^{l*} \cdot \nabla C^*(r^*, \theta^*, t^*)$$

$$\frac{\partial C^*(r^*, \theta^*, t^*)}{\partial t^*} = \frac{D_{Si}}{D_{Fe}} \frac{1}{Pe} \nabla^2 C^*(r^*, \theta^*, t^*) - \mathbf{v}^{l*} \cdot \nabla C^*(r^*, \theta^*, t^*)$$

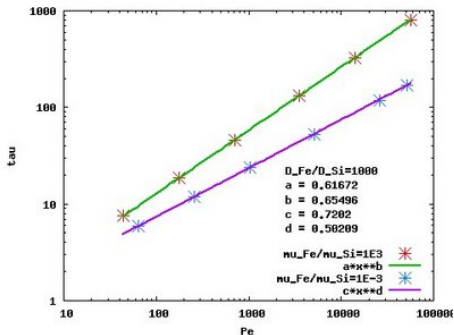
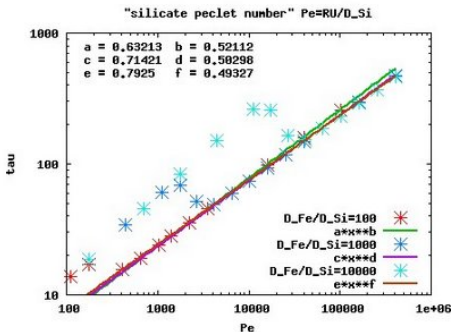
- $Pe = \frac{UR}{D_{Si}}$

# Experiment nb° 1

$$C(t) \sim e^{-\frac{t}{\tau}}$$

two regimes

- $\mu_{Si} > \mu_{Fe} \rightarrow \tau \sim \sqrt{Pe}$
- $\mu_{Si} < \mu_{Fe} \rightarrow \tau \sim Pe^{2/3}$





# Future work

- Investigating of experiment with constant  $K$ .
  - Turn more runs.
  - Find an analytical solution for evolution of concentration.
  - Describe an influence of crucial parameters.
- Experiment with  $K = K(\text{depth})$
- Investigation of the shape of droplets

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Thank you.